



# Infrared Radiative Simulated SSTskin Through Aerosol-Burdened Atmosphere

21st GHRSST Science Team Meeting

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# PART ONE Introduction and Data

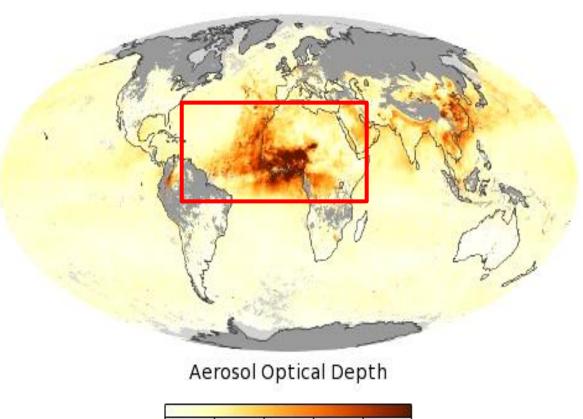


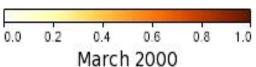
## Significance of this study

- The accuracy of the Sea Surface Temperature (SST) derived from satellite measurements is one of the key factors of climate research and prediction.
- The SST accuracy requirements for climate research are very stringent: ~0.1K.
- But dense tropospheric aerosol concentrations in the atmosphere significantly increase infrared signal attenuation and prevent the retrieval of accurate satellite SSTs.
- Therefore, it is important to quantify the errors and uncertainties of SST and obtain accurate satellite derived SST.



## **Aerosol fields**





**Areas of Interest:** 

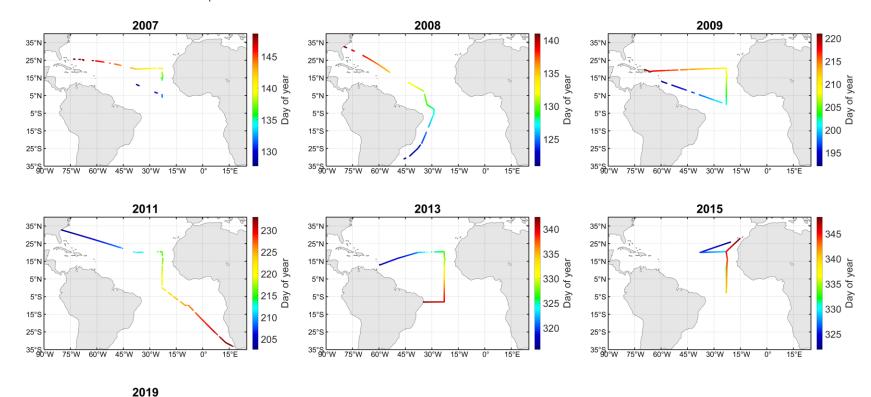
Latitude: 20°S to 35°N

Longitude: 100°E to 100.0°W

Period: From 2006 to 2015

From NASA Earth Observations (NEO): https://earthobservatory.nasa.gov/GlobalMaps/view.php?d1=MODAL2\_M\_AER\_OD





# Aerosols and Ocean Science Expeditions (AEROSE) tracks

Color indicates the days since departure

Nalli, N.R., Joseph, E., Morris, V.R., Barnet, C.D., Wolf, W.W., Wolfe, D., Minnett, P.J., Szczodrak, M., Izaguirre, M.A., Lumpkin, R., Xie, H., Smirnov, A., King, T.S., & Wei, J. (2011). Multiyear Observations of the Tropical Atlantic Atmosphere: Multidisciplinary Applications of the NOAA Aerosols and Ocean Science Expeditions. Bulletin of the American Meteorological Society, 92, 765-789

25°N

15°N

5°N

5°S

35°S0°W 75°W 60°W 45°W 30°W 15°W



#### **Shipboard SST dataset**

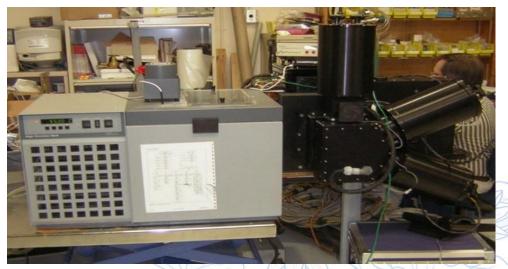
The M-AERI is an accurate, self-calibrating, Fourier transform IR spectroradiometer that measures emission spectra from the sea and atmosphere (Minnett et al. 2001).



NOAA Ship R.H.B at Florida. Mar 2 2018



M-AERI onboard the Ronald H. Brown.



M-AERI is calibrated in the laboratory before and after each deployment using an external validation procedure.



#### **Shipboard Radiosonde dataset**



Table 1. Details of the AEROSE cruises used in this study.

CRUISES	NUMBER OF	START	END	DAYS OF
	RADIOSONDES			DATA
2007 RHB	96	2007-05-07	2007-05-28	22
2008 RHB	74	2008-04-29	2008-05-19	21
2009 RHB	78	2009-07-11	2009-08-11	31
2011 RHB	102	2011-07-21	2011-08-20	31
2013 RHB	111	2013-11-11	2013-12-08	28
2015 Alliance	92	2015-11-17	2015-12-14	28
2019 RHB	97	2019-02-24	2019-03-29	34
Total	650	2007-05-07	2019-03-29	195

RHB: NOAA Ship Ronald H. Brown.

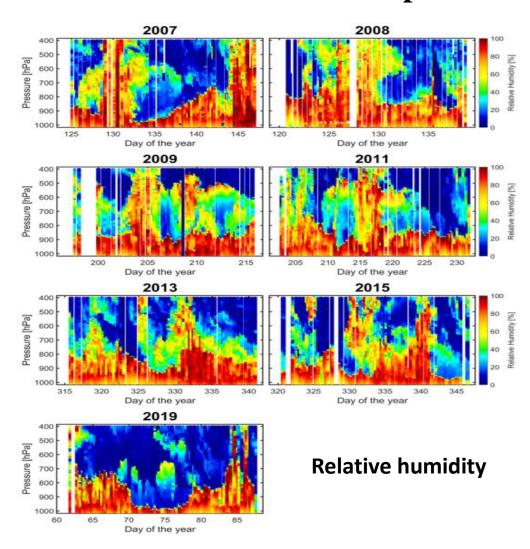
Alliance: North Atlantic Treaty Organization (NATO) R/V Alliance.

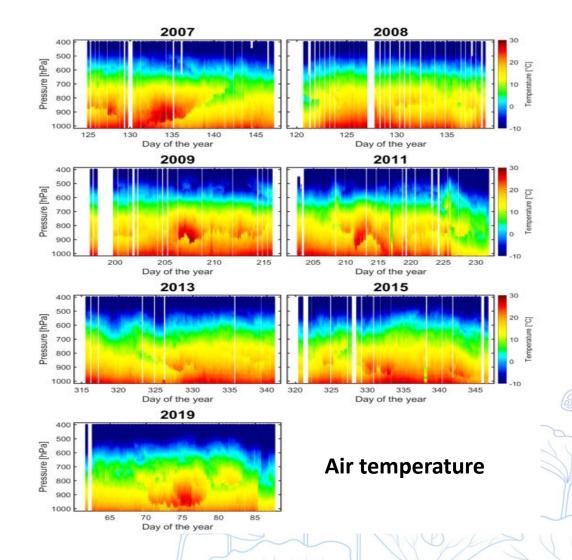
Launching RS92 radiosonde 30 min prior to satellite sounder overpass. (Nalli et al. 2011)





#### **Shipboard Radiosonde measurements**



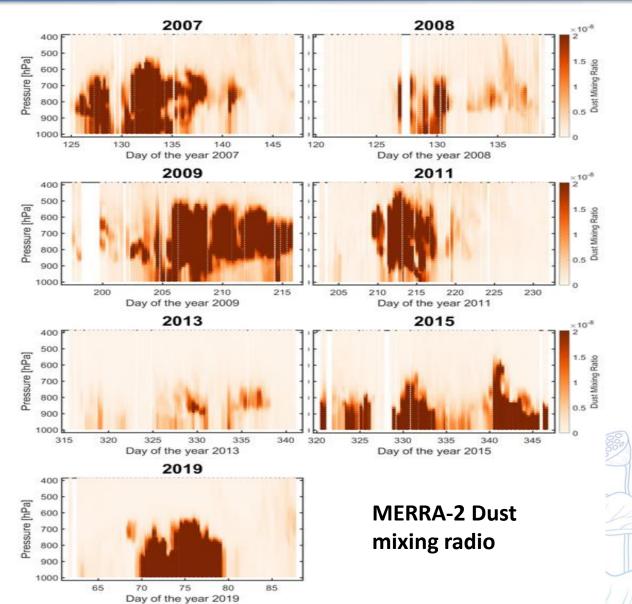




### **MERRA-2** reanalysis value

NASA Modern-Era Retrospective analysis for Research and Applications, Version 2 (MERRA-2) three-dimensional aerosol dust concentrations

Gelaro, Ronald, et al. "The modern-era retrospective analysis for research and applications, version 2 (MERRA-2)." Journal of Climate 30.14 (2017): 5419-5454.





#### **RTTOV**

- RTTOV (Radiative Transfer for TOVS) is a very fast radiative transfer model for measurements of satellite radiometers (Saunders et al. 2018).
- Brightness temperature simulations for TERRA MODIS infrared channels 20 ( $\lambda$ =3.8 µm), 29 ( $\lambda$ =8.9 µm), 31 ( $\lambda$ =11 µm) and 32 ( $\lambda$ =12 µm) have been performed with the RTTOV.
- The brightness temperatures are converted to  $SST_{skin}$  according to MODIS modified nonlinear SST algorithm (NLSST; Walton et al. (1998); Kilpatrick et al. (2015)).
- The aerosol dust-induced  $SST_{skin}$  error described in this study is defined as the aerosol-contaminated SSTskin minus the clear-sky derived  $SST_{skin}$ .
- Input: M-AERI  $SST_{skin}$ , MERRA-2 aerosol dust concentration, Radiosonde Air temperature and Relative humidity

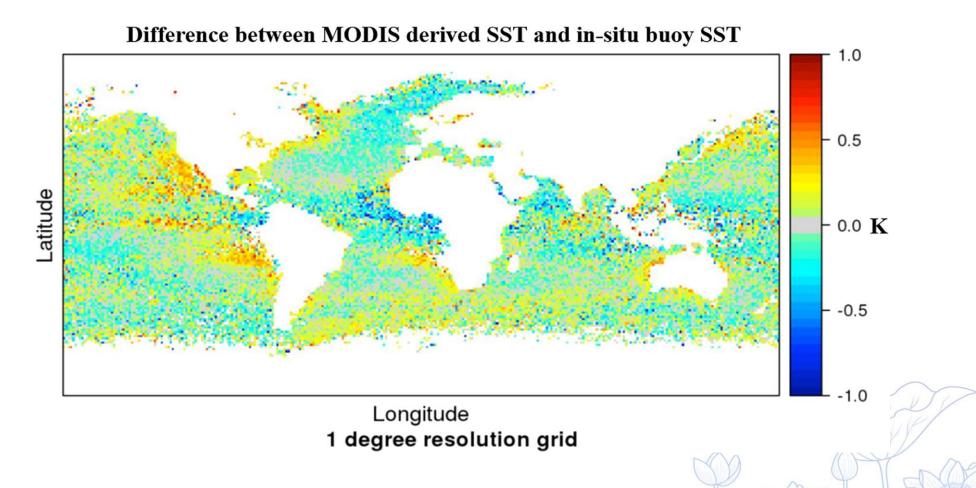
Walton, C. C., et al. "The development and operational application of nonlinear algorithms for the measurement of sea surface temperatures with the NOAA polar-orbiting environmental satellites." Journal of Geophysical Research: Oceans 103.C12 (1998): 27999-28012.

Kilpatrick, K. A., et al. "A decade of sea surface temperature from MODIS." Remote Sensing of Environment 165 (2015): 27-41.



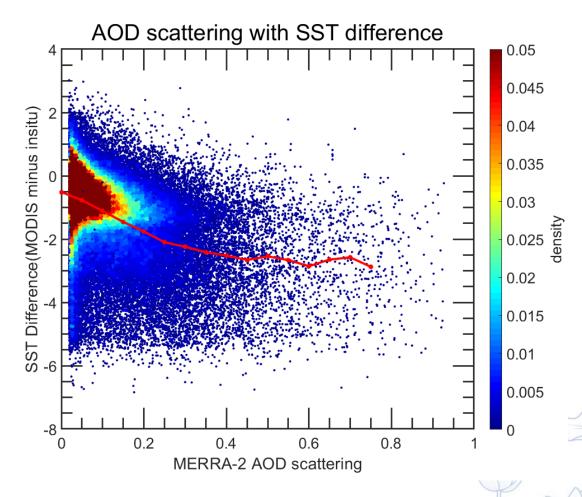
# PART TWO MODIS SST<sub>skin</sub> Validation





Distribution of Terra MODIS SST matchups with in situ drifting buoy measurements (Minnett, 2016).





Scatter plot of SST difference with MERRA-2 Dust Scattering AOD, the red line is the average difference of specific intervals. The SST difference is increasing with AOD.



# PART THREE

RTTOV simulation with AEROSE data



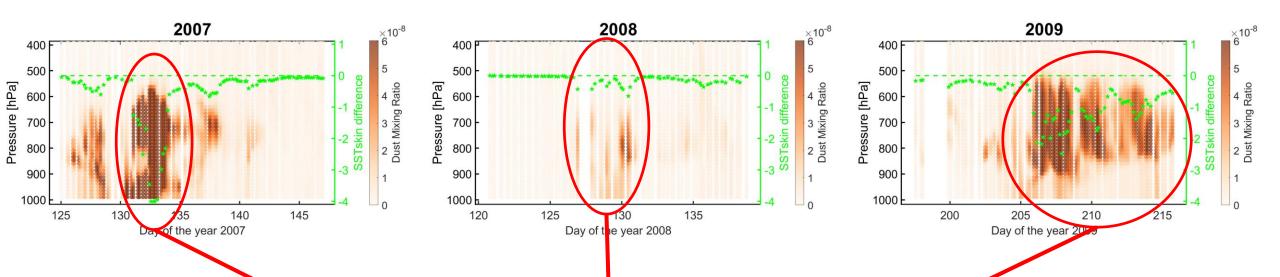


Figure. MERRA-2 aerosol profile along 2007, 2008 and 2009 AEROSE cruises, the color indicates the dust concentrations and the blue dots indicate the simulated SST<sub>skin</sub> difference due to dust aerosol, we use radiosonde profile as input for RTTOV.

SST error is related to dust aerosol concentration and altitude.



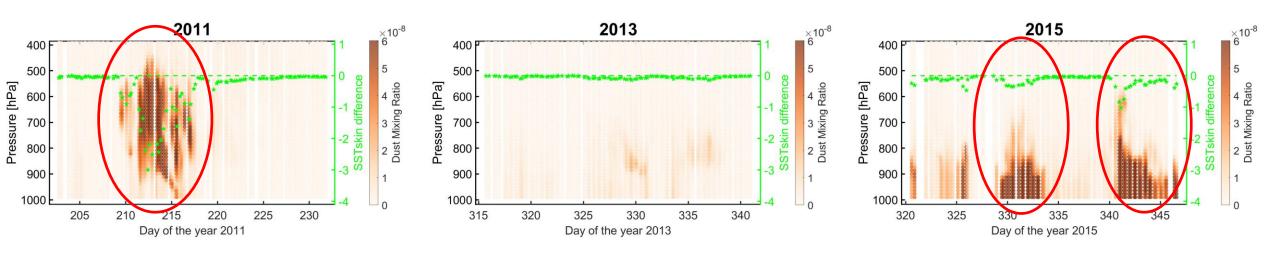
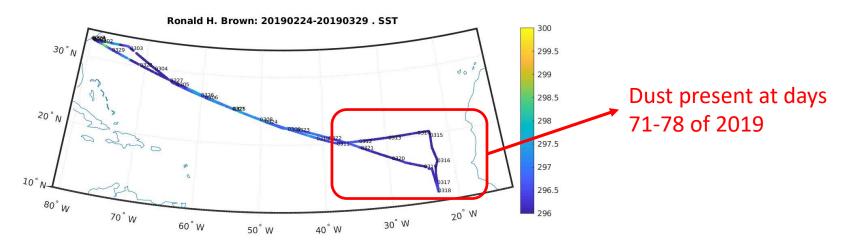
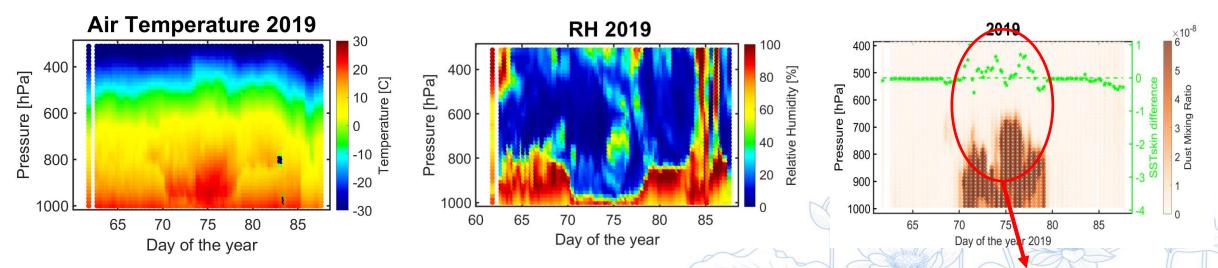


Figure. MERRA-2 aerosol profile along 2011, 2013 and 2015 AEROSE cruises, the color indicates the dust concentrations and the blue dots indicate the simulated SST<sub>skin</sub> difference due to dust aerosol, we use radiosonde profile as input for RTTOV.







RTTOV with TERRA MODIS coefficients interesting that 2019 has positive error! Radiosonde as Atmospheric data input, MERRA-2 as dust input.



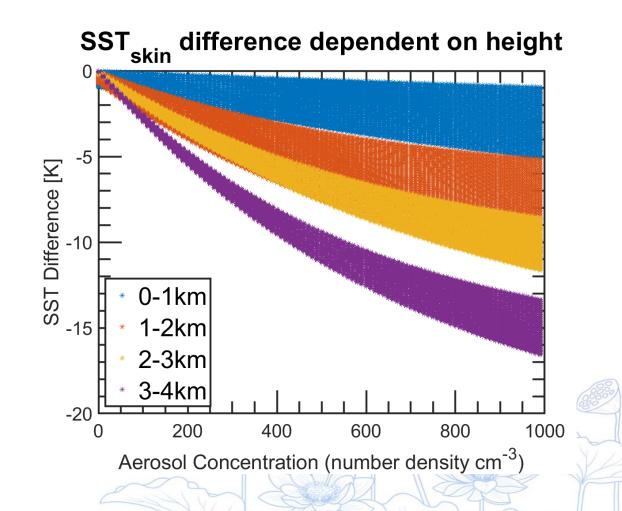
Aerosol vertical distribution effect



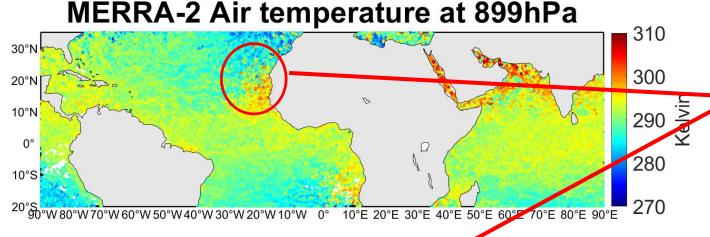
Dust layer altitude and corresponding RTTOV pressure layer

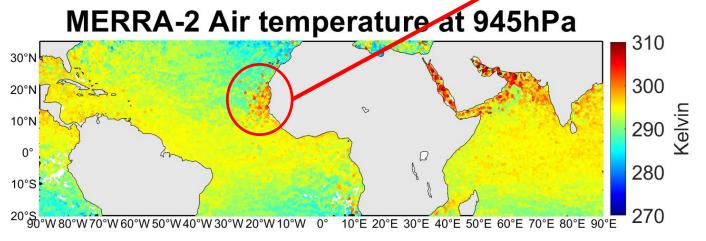
Altitude	Pressure			
0 km - 1 km	922.46 hPa, 957.44 hPa, 985.88 hPa, 1005.43 hPa			
1 km – 2 km	795.09 hPa, 839.95 hPa, 882.8 hPa			
2 km – 3 km	702.73 hPa, 749.12 hPa			
3 km – 4 km	610.60 hPa, 656.43 hPa			

RTTOV simulation of MODIS onboard TERRA satellite



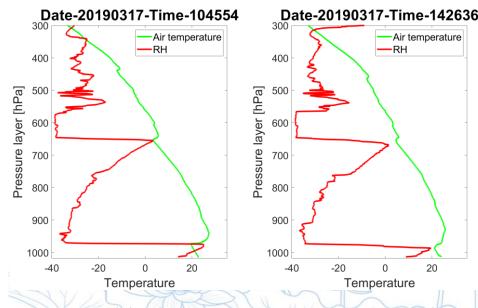




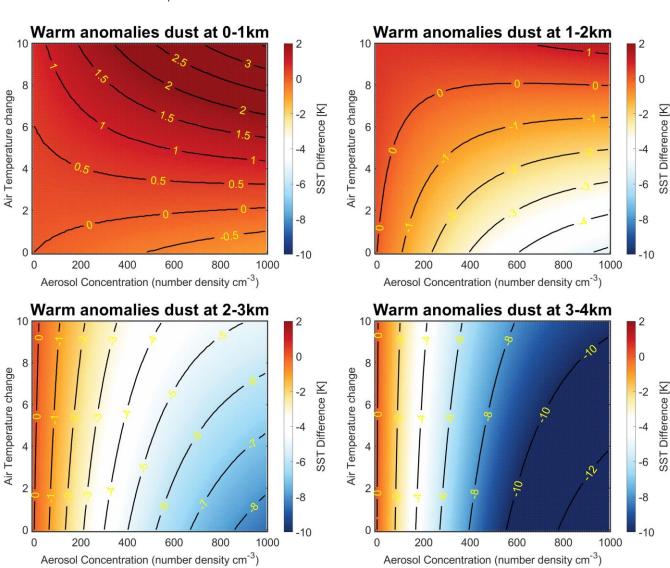


#### **2015-2019 MODIS Terra Matchup Database**

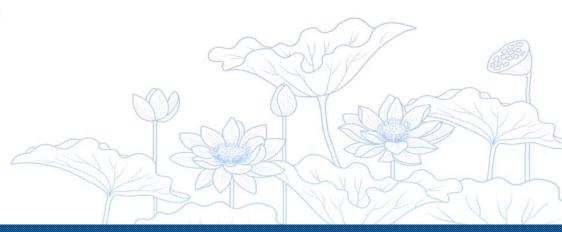
The IR channel brightness temperature is depended on  $\delta T_{sa}$ , which is the difference between the temperature of aerosol layer and surface.







We use RTTOV to simulate the aerosol vertical distribution effects on  $SST_{skin}$ . Vertical distribution of aerosols has a significant impact on  $SST_{skin}$  retrieval. The temperature contrast between SST and the dust layer is related to the retrieved  $SST_{skin}$  error.





# PART FIVE

# Conclusions & Future Work



#### **Conclusions:**

- AEROSE radiosondes and M-AERI provide useful input for RTTOV.
- Temperature changes caused by dust direct radiative effect may result in positive or negative SST<sub>skin</sub> retrieval errors. High concentration dust aerosols in the lower atmosphere warm air temperature, and let the errors positive; dust aerosols in high altitude cause more negative error.
- Reanalysis data, such as those from MERRA-2 and ECMWF ERA-5 (not shown here), can provide vertical aerosol and air temperature data and be used to reduce the satellite retrieved  $SST_{skin}$  errors.





#### **Future Work:**

- The vertical distribution of aerosol influences the accuracy of infrared-derived SST<sub>skin</sub>. RTTOV simulated dust layer effects at different heights have an impact on the satellite SST retrieval.
   CALIPSO provides information about the vertical distribution of aerosol layers, it can be used to derive different coefficients in the correction algorithms.
- The impact of different kinds of aerosol layers should be further explored.
- Such approaches as developed here can be applied to other well-calibrated infrared satellite radiometers such as Visible Infrared Imaging Radiometer Suite (VIIRS) on the Suomi-NPP and NOAA-20, SLSTR on Copernicus Sentinel-3 A/B satellites, Advanced Baseline Imager (ABI) onboard GOES series satellite, and others in the future.



# Acknowledgements





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# THANK YOU!

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