



Seasonal Patterns of SST Diurnal Variation over the Tropical Warm Pool

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2 Data & Method

1 Introduction

- A better understanding of SST diurnal variation (DV) events is essential to better represent the air-sea interaction in weather/climate models (e.g. Clayson and Bogdanoff, [2013]).
- Tropical Warm Pool (TWP, 90°E-170°E, 25°S-15°N) region is considered to be an excellent area to conduct DV studies because of its: (1) globally highest annual average SST, (2) calm winds, (3) strong solar shortwave insolation (SSI), and (4) frequent, large-amplitude DV events.
- Detailed description of the SST DV seasonal patterns using long-term satellite observations is needed in

Major data sets (temporal coverage for five years: Jan. 2010 – Dec. 2014):

- Satellite SST data: Bureau of Meteorology produced daily, 0.02° resolution, version 2 ("fv02") High-Resolution Picture Transmission (HRPT) Advanced Very High Resolution Radiometer (AVHRR) SSTs from NOAA-19, as a contribution to the Integrated Marine Observing System (IMOS; <u>www.imos.org.au</u>); SST field: experimental hybrid SST_Day_Night (Griffin et al. [2016]).
- Drifting buoy SST data: the In-situ SST Quality Monitor (iQuam; www.star.nesdis.noaa.gov/sod/sst/iquam/).
- Hourly wind speed, SSI, and latent heat flux (LH) data: the Bureau's hourly Australian Community Climate and Earth System Simulator-Regional (ACCESS-R) 24-hour forecasts (Puri et al., [2013]). Spatial resolution: 0.375° (0.11°) before (after) 17th Apr. 2013. **Quality control and Method:**
- order to better facilitate DV modelling and to better understand DV features over the TWP region.
- Besides, validation of the new satellite SST product produced by the Bureau of Meteorology is considered necessary before the SST data can be used with confidence in DV studies. In-situ validation against drifting buoy SST data is conducted.
- Only the highest quality level AVHRR SSTs (QL=5) and buoy SSTs (QL=5) are retained;
- SST DV amplitude (dSST) = local daytime SST (~14 local solar time, LST) local night-time SST (~2 LST).

3 In-situ Validation



Fig.1 In-situ validation results: (a) density plot of the **daytime** (~14 LST) collocations within each 0.2 K \times 0.2 K bin for winds > 6 m/s; (b) same as (a) but for **night-time** (~2 LST) collocations for winds > 2 m/s; Please note the Y axes in (a) and (b) are the differences between the satellite and in-situ SSTs. (c) daytime collocation numbers, median bias (solid line), 25th/75th percentiles of the biases (below/above dashed line) over each in-situ SST 1 K range; (d)

4 Seasonal Patterns of SST DV Events



Fig.2 Spatial distributions of the five-year average (a) SST DV, (b) wind speed, (c) SSI, and (d) LH flux. The blank areas in (a) are due to lack of data as the IMOS AVHRR SST data set is derived only from direct broadcast (HRPT) data received at Australian and Antarctic ground stations.



- largest five-year average SST DV The amplitudes can reach 1–2 K, mainly observed close to the equatorial Indonesian islands and Papua New Guinea (PNG) (Fig. 2a), where low wind speeds dominate (Fig. 2b).
- DV events have average amplitudes of around 0.5 K to the north of the north-western Australian coast (NWAC), due to both low winds and high SSI over this region.
- Double-peak seasonal pattern is observed for most of the study domain: maximum DV in Feb - Mar and Oct - Nov; minimum DV in Jun. Strong seasonal pattern is seen in area 1 (red box) due to its relatively high latitudes in the

same as (c) but for night-time collocations. RSD stands for robust standard deviation, R for correlation coefficient.

5 Sensitivity Tests



• For the very weak (WS24<1 ms⁻¹) and the very strong daily average wind speeds (WS24>6 ms⁻¹), the effects of morning winds on dSSTs are small (\pm 0.2–0.4 K). However, for the lowmiddle range daily wind speeds (1 0.35 -0.30 ms⁻¹<WS24<4 ms⁻¹), a calm or very 0.25 windy morning can affect the DV 0.20 -0.15 amplitude up to \pm **0.4-0.6** K. 0.10 When the winds are $< 1 \text{ ms}^{-1}$, the DV 0.05 -0.00 amplitudes are large regardless of the SSI level. The most significant effects $(\pm 0.5 \text{ K})$ that different SSI levels can have are found when the wind speeds are between 1–5 ms⁻¹. Over a fixed wind speed and SSI level, low (high) LH conditions very slightly enhance (reduce) the DV amplitude by **no more than** \pm **0.2** K.

0.19

better, as

Κ

90°E

IMOS

- Double-peak (double-trough) pattern is observed for DV and SSI (wind speed and LH).
- The maximum monthly average DV of **0.33 K** occurs in March. The minimum DV amplitude is found in June. In October, a second peak DV amplitude can

- study domain.
- Weak seasonal pattern is observed in area 2

(blue box) but with year-long scattered large DV

events.



Fig.5 Statistical results of the pixels with different DV amplitudes over the whole study domain and time period. The distribution of the DV amplitudes are shown in (a). The percentages of DV > 1 K, > 2K, and > 3 K against all pixels that have available DV amplitudes in each month are shown in (b).

• The most frequently seen DV amplitudes are within the 0–0.2 K range, which is different from previous studies (e.g. Clayson and Weitlich [2007] and Zhang et al. [2016]) in which the peak DV is in the 0.2–0.4 K range. Reasons: (1) hourly SST data are used in the above two papers, yielding full diurnal cycles and

Fig.6 (a) and (b): Sensitivity test of DV amplitudes to the WS0714 (average wind speed of local morning and early afternoon from 7-14 LST). WS24 means the daily average wind speeds. (c) and (d): Sensitivity of DV amplitudes to SSI levels against different wind speeds. (e) and (f): Sensitivity of DV amplitudes to different LH levels under fixed wind speeds and SSI condition (230 $Wm^{-2} < SSI < 280 Wms^{-2}$). In a, c, and e panels, the sizes of the squares are proportional to the counts of the pixels under that specific condition. The pixel count of the maximum square in a (c, e) is 15238 (7336, 5064) and of the minimum square is 124 (117,120). In b, the max (min) dSST within each 1 ms⁻¹ WS24 range basically corresponds to the weakest (strongest) WS0714 condition within that WS24 range. In d and f panels, the max (min) dSST basically corresponds to the strongest (weakest) SSI and the weakest (strongest) LH condition within that wind speed range. The wind speed condition in SSI and LH sensitivity tests applies to both WS24 and WS0714.

be seen with an amplitude of **0.32 K**.

• The trend for **SSI** is very closely correlated with that of DV. The wind speed trend exhibits good negative correlation with that of DV but for all extreme values, a one-month lag has been found.

reflecting maximum DV amplitudes; (2) daytime IMOS AVHRR SSTs tend to underestimate the true measurements by ~0.2 K.

• The double-peak pattern is revealed in most years for the dSST

> 1 K and > 2 K conditions.

6 Conclusions

- The Bureau produced fv02 IMOS HRPT AVHRR SST (experimental hybrid SST_Day_Night field) from NOAA-19 are of good quality and are suitable for SST DV studies.
- The double-peak seasonal pattern is observed for most of the study domain with maximum DV events found in Feb Mar and Oct Nov and minimum DV events in Jun.
- Sensitivity tests of DV amplitudes to morning and early afternoon (7-14 LST) wind speeds indicates the need of combining daily average winds and the morning winds together in a DV model that only estimates the daily maximum dSST, especially under the conditions when the daily average wind speed is **low-middle** $(1 \text{ ms}^{-1} < WS24 < 4 \text{ ms}^{-1})$.

References

- Clayson, C. A., and A. S. Bogdanoff (2013), The effect of diurnal sea surface temperature warming on climatological air-sea fluxes, J. Clim., 26(8), 2546-2556, doi:http://dx.doi.org/10.1175/JCLI-D-12-00062.1.
- Clayson, C. A., and D. Weitlich (2007), Variability of Tropical Diurnal Sea Surface Temperature, J. Clim., 20(2), 334-352, doi:http://dx.doi.org/10.1175/JCLI3999.1.
- Griffin, C., H. Beggs and L. Majewski (2016), GHRSST compliant AVHRR SST products over the Australian region, in prep.
- Puri, K., G. Dietachmayer, P. Steinle, M. Dix, L. Rikus, L. Logan, M. Naughton, C. Tingwell, Y. Xiao, and V. Barras (2013), Implementation of the initial ACCESS numerical weather prediction system, Aust. Meteorol. Oceanogr. J., 63, 265-284.
- Zhang, H., H. Beggs, L. Majewski, H. Wang, and A. Kiss (2016), Investigating Sea Surface Temperature Diurnal Variation over the Tropical Warm Pool Using MTSAT-1R Data, *Remote Sens. Environ.*, accepted.