

Exploring Internal Wave signature on remote sensing infrared SST observations.

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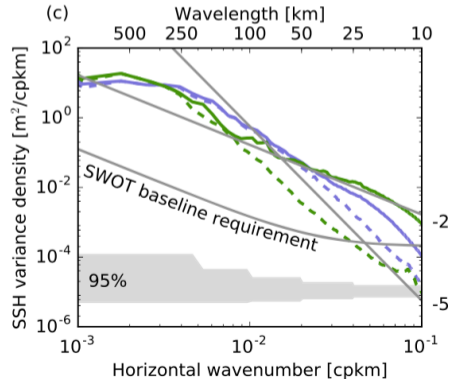
Outline

- 1 Context
- 2 Objective
- 3 Theoretical Background
- 4 Data Sources
- 5 Regions of study
- 6 Preliminary Results
- 7 Summary

Context

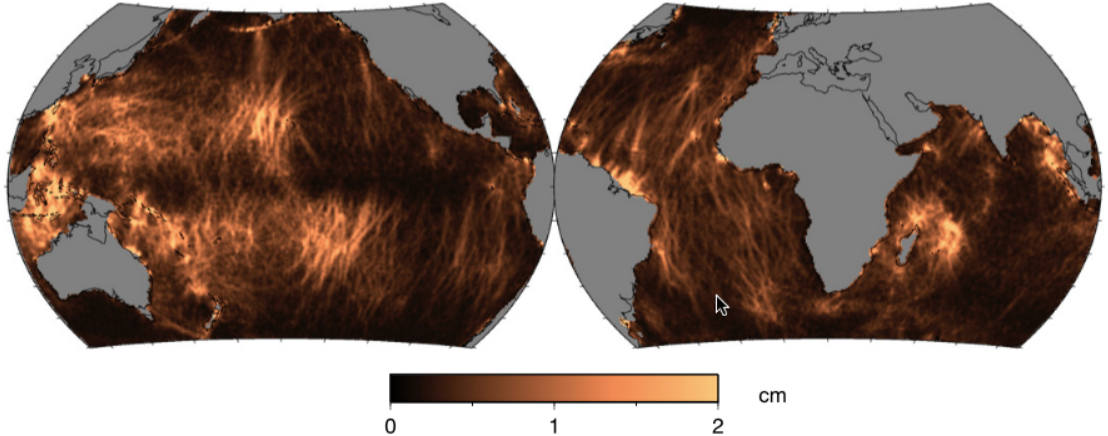
- Internal waves (IW) significantly contribute to sea level variability at scales lower than 100 km

- Fast motions (including IW) will complicate our ability to retrieve ocean currents from altimetry (geostrophy fails)



Context

- Some IW are of tidal origin and stationary
- They can be predicted and removed from altimetric data



- However, a significant fraction of IW are
 - ① From tidal origin but they have lost their phase relationship with respect to astronomical forcing, because of their interactions with slower oceanic turbulence (non stationary internal tides) [Ponte and Klein 2015, Zaron 2016]

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 - 2 From non tidal origin (wind-forced, lee-waves)
- Thus, difficult to predict nowadays

Context

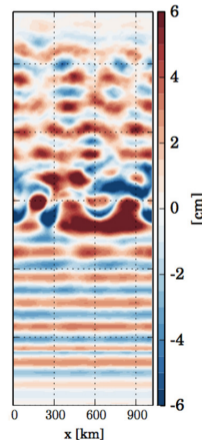
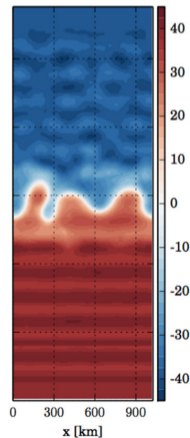
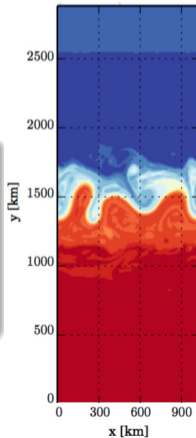
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Context

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- Ponte et al. 2017 proposed in an idealized context a tentative method that relies on the synergy between SST-SSH

2 critical ingredients

- IW filtering capability of the QG framework
- Weakness IW signature on surface tracers(SST)



Objective

- Quantify the IW signature on SST images
- Validate or not the assumption of weak signature IW
- If possible, identify regions with some internal tides signature, strong SST gradients, cloud free regions.

Theoretical Background

- Quantify the IW signature on SST images
 - IW currents (u_w, v_w) are expected to periodically advect SST (T_s) fronts and result in SST fluctuations (T_w):

$$\partial_t T_w = -u_w \partial_x T_s - v_w \partial_y T_s$$

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- Assuming that $u_w = \Re(u_c e^{-iwt})$ where $u_c = u_r + iu_i$ and w is the IW frequency

$$\Re(-iwt_w e^{-iwt}) = -\Re(e^{-iwt} [u_c \partial_x T_s + v_c \partial_y T_s])$$

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- Assuming that $u_w = \Re(u_c e^{-i\omega t})$ where $u_c = u_r + iu_i$ and ω is the IW frequency

$$\Re(-i\omega T_w e^{-i\omega t}) = -\Re(e^{-i\omega t} [u_c \partial_x T_s + v_c \partial_y T_s])$$

- Thus, the amplitude of IW signature on SST can be estimated as:

$$T_w = \frac{u_c \partial_x T_s + v_c \partial_y T_s}{i\omega}$$

● Tidal atlas

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• Tidal atlas

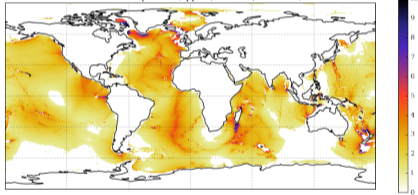
• SST remote sensing

Data Sources

- Atlases of **baroclinic** and **barotropic** sea level (currents):

Baroclinic

M2 current amplitudes (approx, $a > 1\text{cm/s}$, $h > 1000\text{m}$)



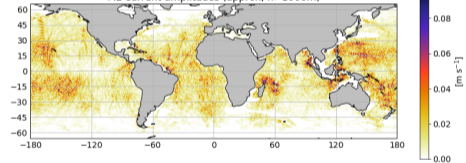
HRET

(High Resolution Empirical Tide) [Zaron]

Spatial resolution = $1/20^\circ$

Barotropic

M2 current amplitudes (approx, $h > 1000\text{m}$)



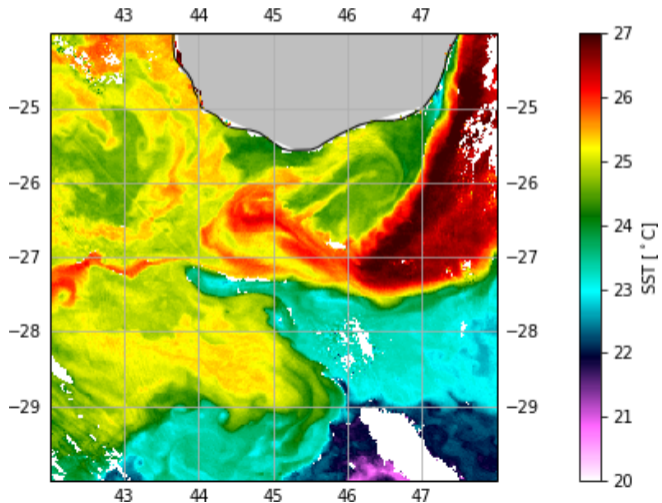
FES

(Finite Element Solution)

Spatial resolution = $1/16^\circ$

Data Sources

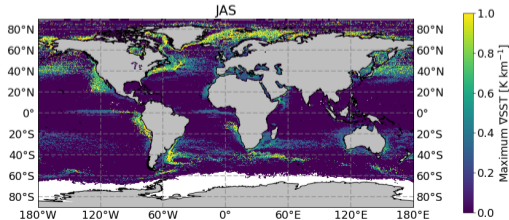
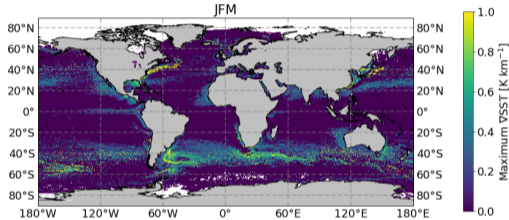
- L2 SST granules from: METOP, VIIRS, MODIS (2014-2016)



L2 SST METOP
06/06/2014
Spatial resolution = 0.02°
(remapped)

Data Sources

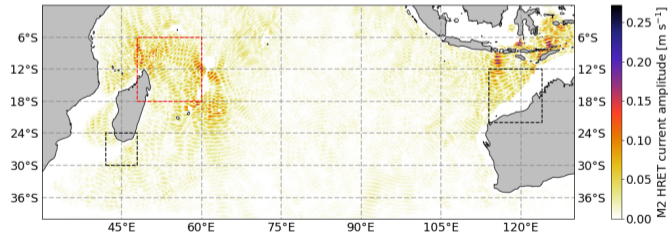
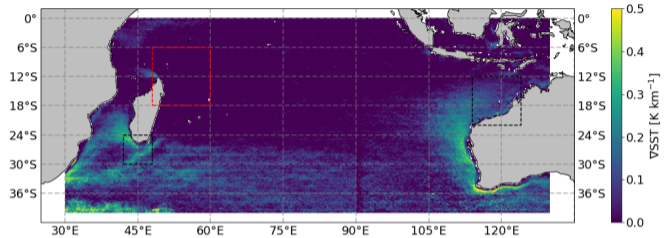
- Climatology of Maximum SST gradient. Courtesy of Peter Cornillon , Graduate School of Oceanography, University of Rhode Island (URI).



- Obtained from the entire (1985-1996) Pathfinder 9 km resolution SST dataset
- Based on the automated procedure by Cayula 1991
- Spatial resolution 9 km

Regions of study

- Strong thermal SST gradients
- Significant signature of IW



Preliminary Results

- Signature of IW on SST:

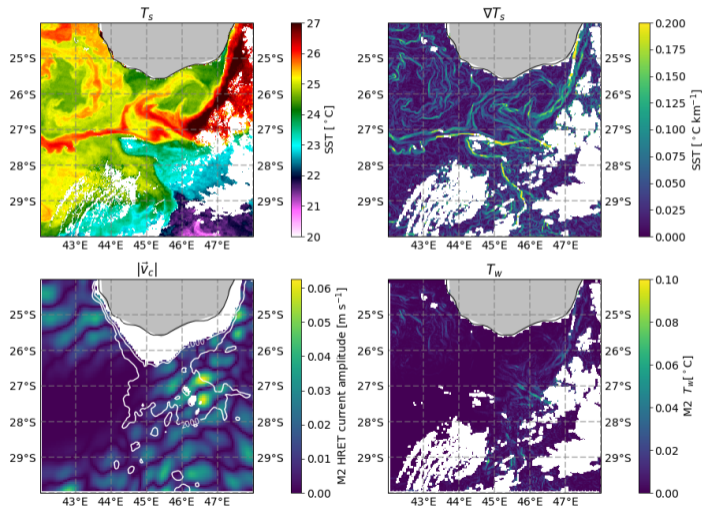
$$T_w = \frac{u_c \partial_x T_s + v_c \partial_y T_s}{i\omega}$$

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M2 Baroclinic ($w = 1.932\text{cpd}$)

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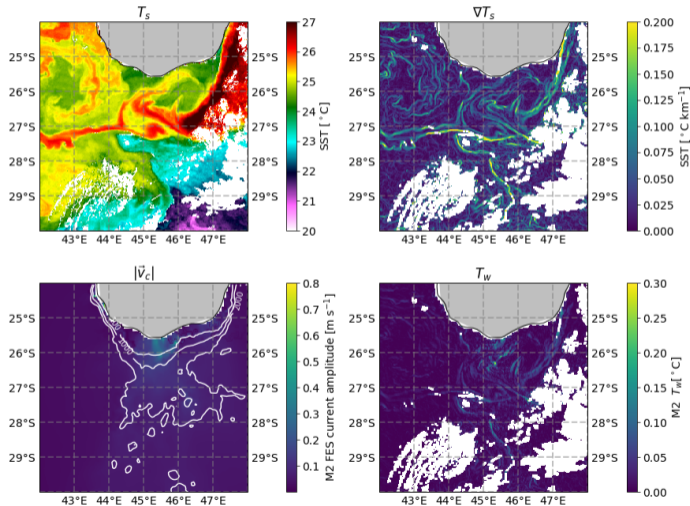
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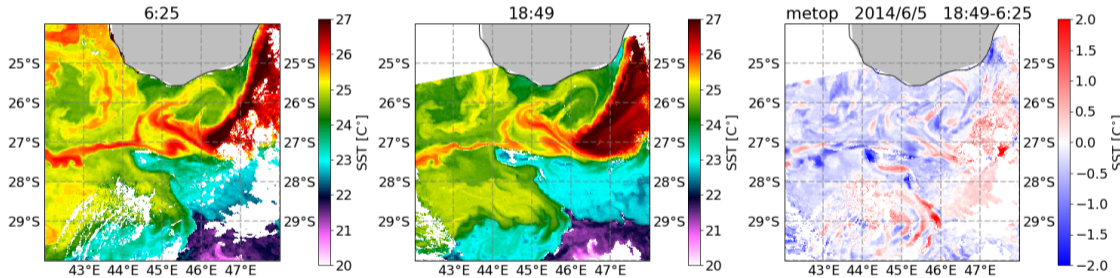
M2 Barotropic ($w = 1.932\text{cpd}$)

$$T_w = \frac{u_c \partial_x T_s + v_c \partial_y T_s}{i\omega}$$



Preliminary Results

- Temporal gradient between consecutive passes



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- This **signature** is related to the **advection** of **fine scale SST fronts** (fine-scale and confined to the front location). It is proportional to the product of the SST frontal gradient and tidal currents.

Summary

- We present here a first attempt to **quantify** the signature of **IW** on **SST observations** provided by **orbiting satellites**.
- This **signature** is related to the **advection** of **fine scale SST fronts** (fine-scale and confined to the front location). It is proportional to the product of the SST frontal gradient and tidal currents.
- **SST fluctuations** due to **barotropic** tidal motions are lower than **0.3 °C**.
- **SST fluctuations** due to **internal** tides motions are lower than **0.1 °C**.

- Such **signatures** appear to be **weak** compared to **SST variations over short** (tidal) **temporal windows** associated to: mesoscale/submesoscale; diurnal cycles.

- Such **signatures** appear to be **weak** compared to **SST variations over short** (tidal) **temporal windows** associated to: mesoscale/submesoscale; diurnal cycles.
- These **amplitudes** may be **compared** to the **absolute accuracy** of **SST products** ($\simeq 0.3$ K) for the M2 barotropic and lower for M2 internal tides [Ocarroll et al. 2012, Wu et al. 2017] and to the **instrument pix noise** (presumably weaker than the former).

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